Soil Development

Soils Sustain Life

Right off the Bat!!

- A soil is not a geologic deposit, it is a product of in situ weathering.
- Engineers, archaeologists, and others have learned that anything that is not rock on the Earth's surface is soil.

The way a soil is defined depends largely on the way you view it, eg. engineers, archaeologists, agronomists, ect.

Our definition of Soil

A soil is a natural body consisting of layers (horizons) of mineral and/or organic constituents of variable thickness, which differ from the parent materials in their morphological, physical, chemical, and mineralogical properties and their biological characteristics.

Soil forming factors

Hans Jenny (1941) Cl, o, r, p, t ► cl, climate ▶ o, biotic influence r, topographic relief ▶ p, parent material ▶t, time

Jenny's individual soil functions

- s or S = f (<u>cl</u>, o, r, p, t)
- ► s or S = f (<u>o</u>, cl, r, p, t)
- s or S = f (<u>r</u>, cl, o, p, t)
- Lithofunction
- Chronofunction
- To solve each function, the first factor listed (<u>cl</u>) is allowed to vary while the others remain constant. It is then possible to statistically determine the dependancy of one (or more) soil properties on a single factor

climofunction biofunction topofunction

Soil characteristics

► Horizons

| Horizon | Properties | | |
|---------------|--|--|--|
| 0 | Uppermost horizon; fresh to partly decomposed organic matter. | | |
| А | Mineral soil mixed with humus; dark colored. | | |
| Englished and | Zone of maximum <i>eluviation</i> (leaching) of clay, iron, and aluminum, leaving residual minerals such as quartz in sand and silt sizes. Lighter color than A horizon. | | |
| В | Zone of <i>illuviation</i> (accumulation) of clay, iron, and aluminum compounds from above, with development of distinctive structures such as granular, blocky, or prismatic aggregates. Colors more intense than overlying and underlying horizons. | | |
| C | Relatively unweathered unconsolidated mineral matter below the zone of major biologic activity. May be mechanically fractured, stained by oxides, or loosely recemented by calcium carbonate, gypsum, iron oxide, or more soluble salts. | | |





Horizons

Distinctive weathered zones that are roughly parallel to the land surface and are the product of weathering processes. The Master horizons are O, A, E, B, C, and R.
 Subhorizons

Horizon examples

Soil Profile



Leaf litter

Mineral horizon at the surface showing organic matter enrichment

Subsurface horizon showing depletion of organic matter, clay, iron, and aluminium compounds

Subsoil horizon showing enrichment of clay material, iron, aluminum, or organic compounds

Horizons of loosened or unconsolidated material

Hard Bedrock



Subhorizons

TABLE 2.6 Lowercase Letter Symbols to Designate Subordinate Distinctions Within Master Horizons

| Letter | Distinction | Letter | Distinction |
|--------|--|--------|---|
| a | Organic matter, highly decomposed | n | Accumulation of sodium |
| b | Buried soil horizon | о | Accumulation of Fe and Al oxides |
| c | Concretions or nodules | р | Plowing or other disturbance |
| d | Dense unconsolidated materials | ģ | Accumulation of silica |
| e | Organic matter, intermediate decomposition | r | Weathered or soft bedrock |
| f | Frozen soil | S | Illuvial accumulation of O.M. ^a and Fe and Al oxides |
| g | Strong gleving (mottling) | SS | Slickensides (shiny clay wedges) |
| h | Illuvial accumulation of organic matter | t | Accumulation of silicate clays |
| i | Organic matter, slightly decomposed | v | Plinthite (high iron, red material) |
| i lot | Jarosite (yellow sulfur mineral) | w | Distinctive color or structure |
| ii . | Cryoturbation (frost churning) | x | Fragipan (high bulk density, brittle) |
| k | Accumulation of carbonates | v | Accumulation of gypsum |
| m | Cementation or induration | z | Accumulation of soluble salts |

*O.M. = organic matter.

Subhorizon characteristics

Surface horizons = epipedons

| Mollic (A) | Thick, dark-colored, high base saturation, strong structure | | |
|---------------------|--|--|--|
| Umbric (A) | Same as mollic except low base saturation | | |
| Ochric (A) | Too light-colored, low organic content or thin to be Mollic; may be hard and massive when dry | | |
| Melanic (A) | Thick, black, high in organic matter (>6% organic C), common in volcanic ash soils | | |
| Histic (O) | Very high in organic content, wet during some part of year | | |
| Anthropic (A) | Human-modified molliclike horizon, high in available P | | |
| Plaggen (A) | Human-made sodlike horizon created by years of manuring | | |
| Subsurface horizons | | | |
| Argillic (Bt) | Silicate clay accumulation | | |
| Natric (Btn) | Argillic, high in sodium, columnar or prismatic structure | | |
| Spodic (Bh, Bs) | Organic matter, Fe and Al oxides accumulation | | |
| Cambic (Bw, Bg) | Changed or altered by physical movement or by chemical reactions, generally nonilluvial | | |
| Agric (A or B) | Organic and clay accumulation just below plow layer resulting from cultivation | | |
| Oxic (Bo) | Highly weathered, primarily mixture of Fe, Al oxides and non-sticky-type silicate clays | | |
| Duripan (Bqm) | Hardpan, strongly cemented by silica | | |
| Fragipan (Bx) | Brittle pan, usually loamy textured, dense | | |
| Albic (E) | Light-colored, clay and Fe and Al oxides mostly removed | | |
| Calcic (Bk) | Accumulation of CaCO ₃ or CaCO ₃ · MgCO ₃ | | |
| Gypsic (By) | Accumulation of gypsum | | |
| Salic (Bz) | Accumulation of salts | | |
| Kandic (Bt) | Accumulation of low-activity clays | | |
| Petrocalcic (Ckm) | Cemented calcic horizon | | |
| Petrogypsic (Cym) | Cemented gypsic horizon | | |
| Placic (Čsm) | Thin pan cemented with iron alone or with manganese or organic matter | | |
| Sombric (Bh) | Organic matter accumulation | | |
| Sulfuric (Cj) | Highly acid with Jarosite mottles | | |

Transitional major horizons

► AB, BA, OE, ect.



Formational factors considerations

- Each certainly influences the end weathering product (soil)
- Climate, organisms, and topographic relief, are likely to change (evolve) over the duration (2 to 20,000 years) of soil formation.

In general, Climate (ppt. and temp.) and Parent material are the most controlling factors in regards to soil formation. (complex combinations.

An example of the importance of climate

- Australian buaxite
- Parent material, 90% sandstone, 10% kaolinite
- Climate, Hot monsoons, with approx. 1500mm/yr ppt., plus time produces an abundant source of buaxite from and unlikely parent material.
- The end weathering product (soil profile) has leached the majority, 85% of the quartz to 8 m, the upper 5 m consists of hydrated AIOH (oxides) with iron stains.

Parent Material types



The basic processes of soil formation

Transformations
 Translocations
 Additions
 Losses

Transformations

Chemical and mineralogical changes or precipitation ► Hydrolosis Neoformation Decomposition of organic residue and synthesis of organic

acids, humus



Translocations

The movement (laterally or vertically) of inorganic or organic materials within a soil profile



Translocation terms

Eluviation, The removal of soil material in suspension from a layer or layers of soil (leaching)

Illuviation: A soil layer or horizon in which material is carried from an overlying layer. (A layer of accumulation).

Additions

- Organic matter
 Precipitation, aqueous, salts/silica (evaporation)
- Deposition

Losses

Leaching
Evapotranspiration
Erosion



The four processes of soil formation

These processes, Transformations, Translocations, Additions, and Losses operate under specific environmental factors and provide a framework to investigate the relationships between soils and the landscapes/ecosystems in which they are active.

Soil Classification

Landscape/Environment, Polypedon, pedon, epipedon

Soil landscape



A "pedon"

Epipedon

Includes the upper portion of the soil that is darkened by organic matter and/or the upper eluvial horizons. ► The epipedon may extend into the Bhorizon if it is darkened by organic material.



Soil Taxonomy

Order 12
Suborder 63
Great Group 319
Sub Group 2, 484
Family 8, 000
Series 19, 000

Soil and Time



Describing a soil

Color
Thickness
Structure
Texture

Structure

As a soil develops, soil peds begin to form shapes or a structure

Granular

A-horizon

B-horizon

Platy

Prismatic

Columnar

Texture

Size of the individual peds

Soils and Geomorphology

The Geography of Soil Development
 Importance:

what, why, WHERE? It is important to understand where developing soil assemblages occur.

Agriculture, engineering, septic tank placement, etc.

Geographic variability and Soils

Small scale Your back yard Medium scale Iowa Large scale The Great Plains

Small scale soil variability

A product of small changes in topography and thickness of parent material and/or the effects of organisms.

Medium scale soil variability

Soil forming factors and individual soil assemblages
Lithosequences
Chronosequences
Toposequences
E.g. inceptisols

The Catena concept

A sequence of different soils formed under similar conditions except for the effect of their relief positions effect the drainage and erosion rates of each different soil.

Catena

Helps relate soils to a particular landscape

Soil associations

A general grouping of individual soils that occur together in a particular landscape.

Soil associations

Soil associations

Large-scale soil variability

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Global Soil Regions

E.g. Oxisols, found in hot, humid, regions Mollisols, found in Semi-arid grasslands Aridisols, found in desert regions

LATERITE

(a)

Wet climate

Thin or absent humus

Thick masses of insoluble iron and aluminum oxides; occasional quartz

Iron-rich clays and aluminum hydroxides

Thin leached zone

Mafic igneous bedrock

Paleosols

A soil that formed on a landscape of the past. –Ruhe 1965
A soil that formed in a previously existing climate and may know be buried.

Paleosols

- Paleosols are <u>ancient</u> or <u>fossil soils</u> preserved in the stratigraphic record
 Paleosols generally occur beneath <u>unconformities</u>
- Paleosols provide good clues for the interpretation of <u>paleoenvironments and</u> <u>paleoclimates</u>

Paleosols

Criteria for recognizing paleosols:

Presence of recognizable soil horizons

- Organic upper horizon
- Subhorizons exhibiting increasing clay and/or reddish color
- Uppermost surface sharply truncated

Chemically different than parent material including gradational mineral content, with unstable minerals decreasing in abundance up-section

- 1. Presence of soil structures (e.g. peds)
- 2. Laterally extensive

3. Root traces, or other disruption of original sedimentary structures by organic activity

Paleosol types

Relict soils, exposed, unearthed, paleosols
 Fossil soils, paleosols buried in a geologic section and now lithified.

Paleosol Controversy

Paleosol, Vs. Environmental Change (Intensity/time/geography)

The Holocene

Soils developed on old surfaces that have underwent significant changes yet do NOT contain characteritics of a previous environment

Diagenesis

A Paleosol

Soil structures: *Root traces*

